# Fluid inclusions of ophicarbonate at the Piemont Zone, Western Alps, as a case of paleo oceanic core complex at slow-spreading ridges

Tomohiro Inukai <sup>1,</sup> \*, Mikiya Kageta <sup>1</sup>, Tatsuhiko Kawamoto <sup>1</sup>, Miki Tasaka <sup>1</sup>, Hajime Taniuchi <sup>1,2</sup>, Kenneth T. Koga <sup>3</sup>, Christian Nicollet <sup>4</sup>, Estelle F. Rose-Koga <sup>3</sup>, Baptiste Debret<sup>5</sup> <sup>1.</sup> Shizuoka University Department of Geosciences, Japan <sup>2.</sup> GSJ, AIST, Japan <sup>3.</sup> Institute des Sciences de la Terre à Orléans, CNRS, Université d'Orléans, France <sup>4.</sup> Université Clermont Auvergne, CNRS, France

- <sup>5.</sup> Institut de physique du globe de Paris, France

# **1. Introduction**

### **Carbonation of serpentinite**

Contribution to the global carbon cycle (e.g. Alt et al., 2013, Lithos) Relation to slow earthquakes (e. g. Okamoto et al., 2021, Commun. Earth Environ.) Carbon capture and storage (e. g. Kelemen et al., 2011, Annu. Rev. Earth Planet. Sci.)

The oceanic core complex, where the mantle is exhumed at the seafloor, forms carbonated serpentinite (ophicarbonate).

### This study

Clarifies the chemical species, salinity and temperature of the fluid that forms ophicarbonate at the oceanic core complex

# **2. Sampling locality** European/Iberia units Oceanic unit Granites (Cenozoic) Mantschal et al. (2010, Lithos)

The Piemont Zone is an oceanic unit containing a paleo oceanic core complex (Manatschal et al., 2010, Lithos). With a series of metamorphic grades, this unit is well suited to study chemical evolution in subduction zones. We sampled from two ophiolites; Chenaillet (almost free of subduction metamorphism) and Lago Nero (metamorphism of green-schist facies).

# 3. Methods

Field emission-scanning electron microscopy (FE-SEM) microstructure observation and chemical analysis of thin sections

Raman microscopy identification of mineral and chemical species in fluid inclusions

Heating and cooling stage measuring salinity and homogenization temperature of fluid inclusions

# 6. Conclusion

Fluid inclusions suggest that the calcite core and matrix were formed during seawater interaction.

The temperature of carbonate formation can be related to subduction depths.

We will analyze samples from ophicarbonate associated with higher-grade metamorphic rocks.









### **Chenaillet ophicarbonate**

Serpentinite breccias are filled with calcite **matrix** and **veins**. Serpentinite is locally replaced by calcite **cores**.



Lizardite contains lower concentrations of AI than chrysotile.





Andradite (Ca<sub>3</sub>(Fe<sup>3+</sup>,Ti)<sub>2</sub>SiO<sub>3</sub>O<sub>12</sub>) was locally observed in the calcite matrix. The presence of **Fe<sup>3+</sup>** in andradite suggests that the fluids were highly oxidized (Cannaò et al., 2020 Chem. Geol.).



Lago Nero ophicarbonate Branching calcite veins develop in the serpentinite.







Calcite core (Chenaillet)





Calcite vein (Lago Nero)





## Lago Nero

Differences in the homogenization temperatures of fluid inclusions between **Chenaillet and Lago Nero** samples may represent differences in the subduction depths.



Chenaillet **Oxygen isotope thermometry** assuming **seawater** as the fluid estimated calcite cores at 130-160°C and matrix calcites at 155-**170 °C** (Lafay et al., 2017, Lithos). Similar homogenization temperatures indicate that these calcites were formed from seawater.

Carbon isotope results suggested that calcite cores were formed with serpentinite (Lafay et al., 2017). Fluid inclusions with higher salinity suggest that the calcite cores were formed by fluids associated with serpentinization (Debure et al., 2019, Sci. Rep.).

Matrix calcite was precipitated from seawater in the hydrothermal <sup>300</sup> circulation of the seafloor (Lafay et al., 2017). Depression and **boiling** of seawater resulted in a **wide range of salinities** (Kelly & Delaney, 1987, Earth Planet. Sci. Lett.; Lafay et al., 2017).



